

Numerical modeling of cracking process in partially saturated porous media and application to rainfall-induced slope instability analysis

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1 . Background

2 . Phase-field formulations

3 . Numerical Modeling

4 . Conclusions and Perspectives

01

PART ONE

BACKGROUND

Introduction

Landslide

Predisposing factors

- Lithology
- Faults
- Land use

Easily assessed by spatial analysis techniques
(Ayalew et al. 2005)

Triggering factors

- Rainfall
- Snowmelt
- Earthquakes

Difficult to estimate at a regional scale
(Griffiths 2014)

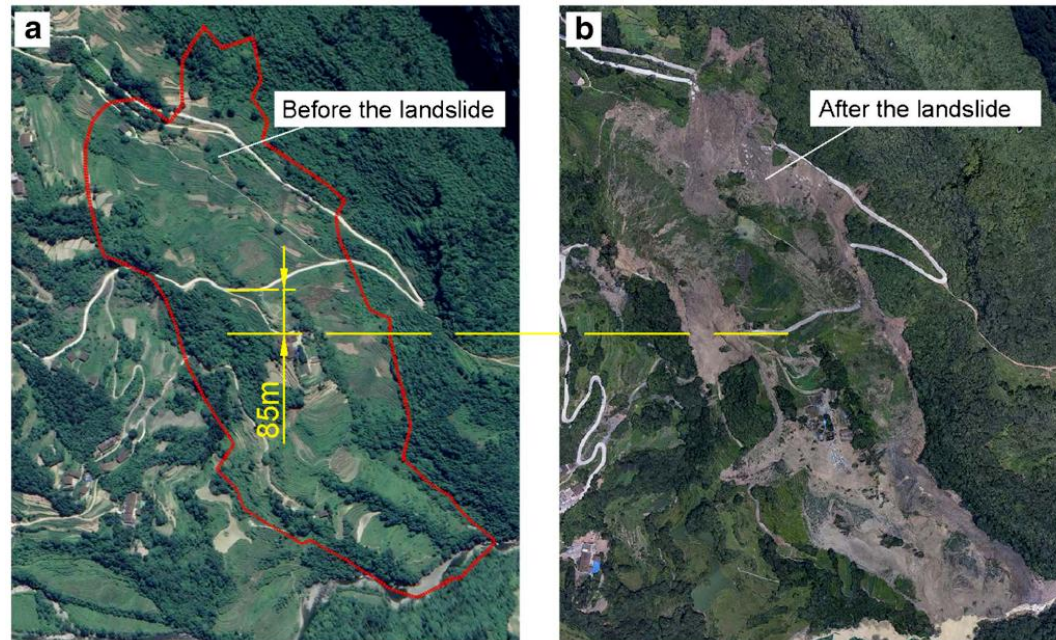


Figure –Zhongbao lanslide at Wulong, China
(CHEN et al. 2021)

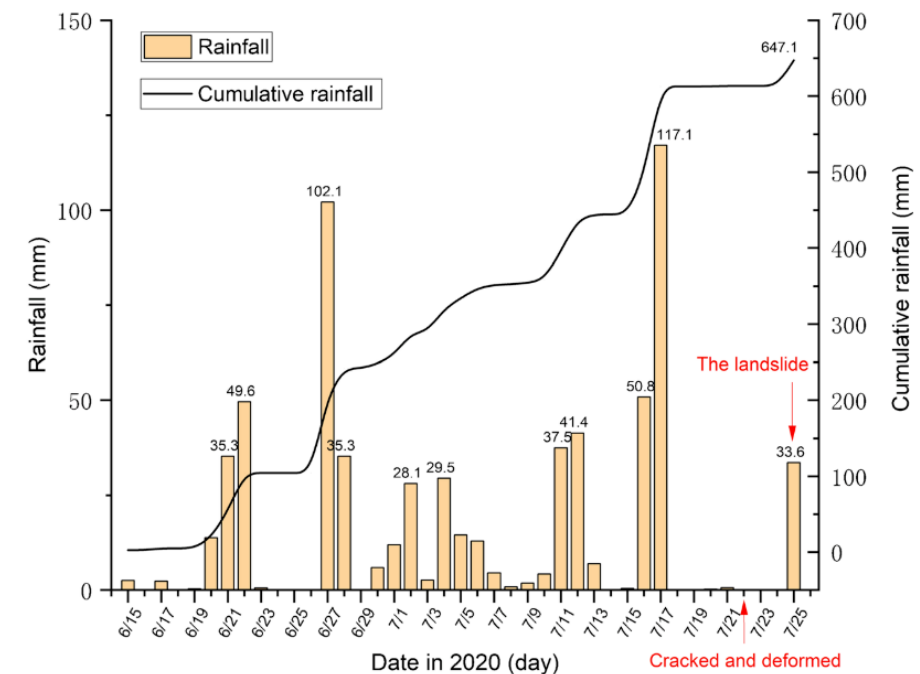
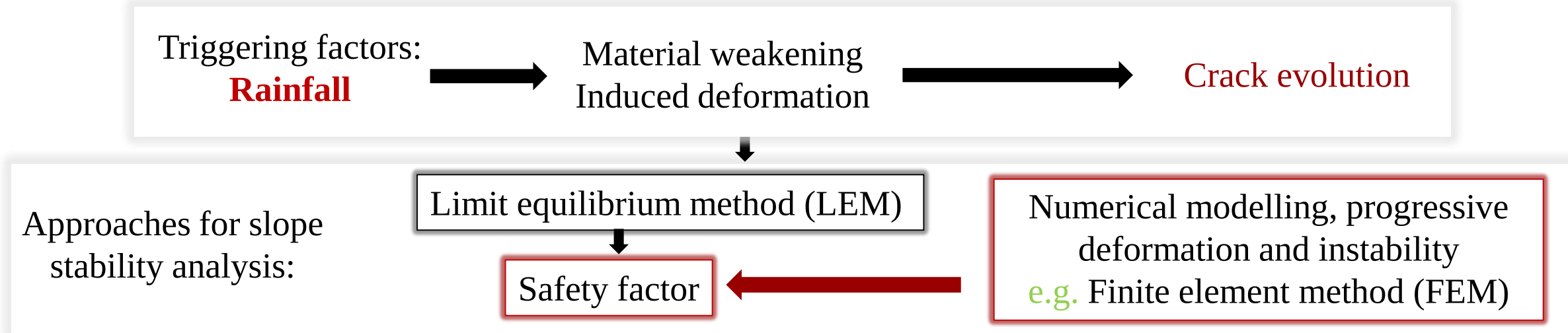


Figure –Rainfall monitoring data
(CHEN et al. 2021)

Objectives



Methodology: Phase-field Method (PFM)

Regularized crack topology:

$$A_{\Gamma} = \int_{\Gamma} dA \cong \int_{\Omega} \gamma_d(d, \nabla d) dV$$

- Predict not only **crack initiation** but also the **crack propagation path**;
- Deal with merging and branching of **multiple cracks**;
- Easy to incorporate the **multi-field physics**

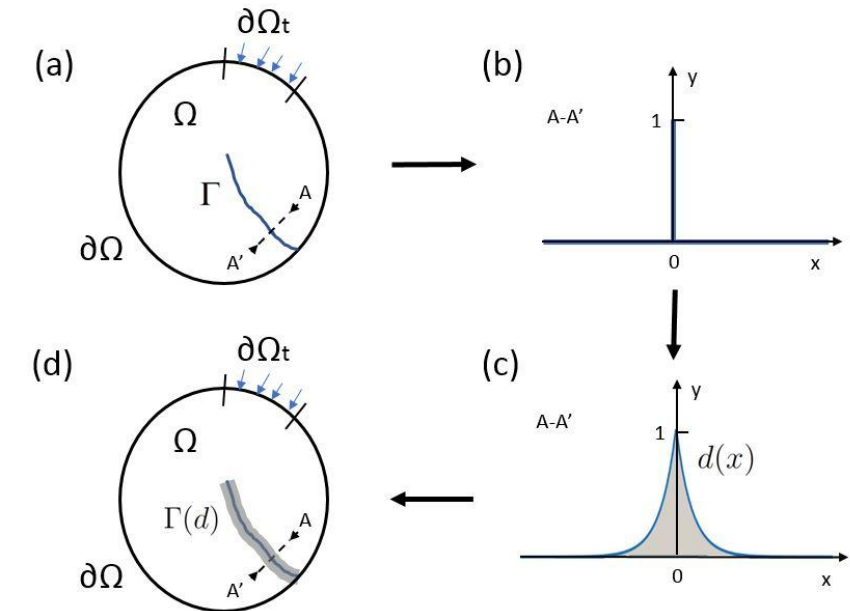


Figure –Regularized crack topology

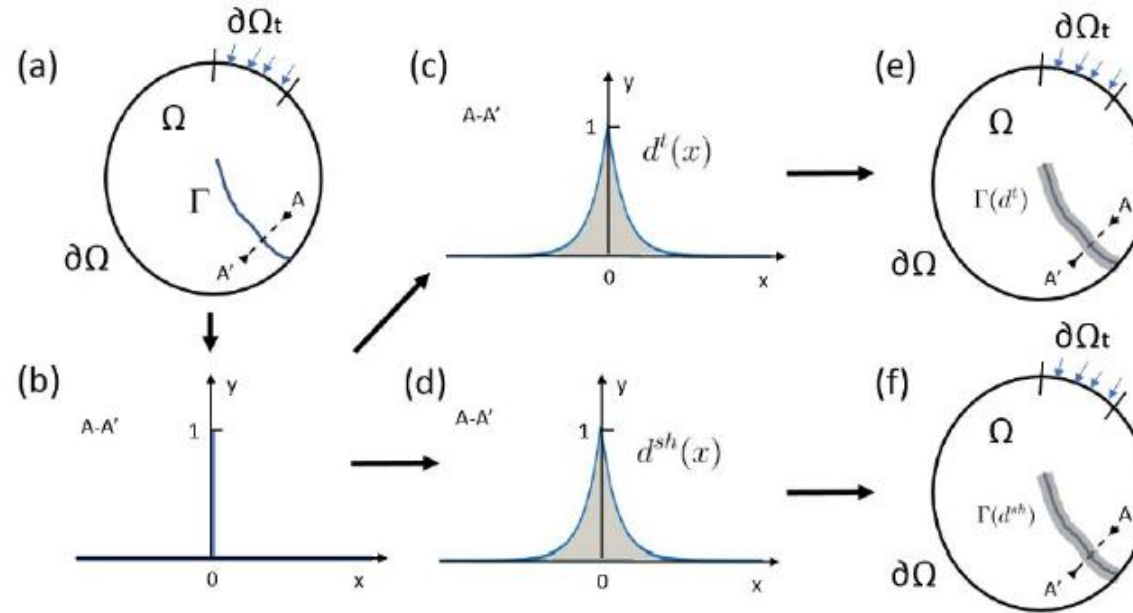
02

PART TWO

PHASE-FIELD FORMULATIONS FOR PARTIALLY SATURATED MEDIA

Regularized crack fields

Two independent variables d^t and d^s to approximate the crack surface area:



The total crack surface density : $\gamma_d(d, \nabla d) = \gamma_d^t(d^t, \nabla d^t) + \gamma_d^s(d^s, \nabla d^s)$

Tensile crack density

$$\frac{(d^t)^2}{2l_d} + \frac{l_d}{2} \nabla d^t \cdot \nabla d^t$$

$$\frac{(d^s)^2}{2l_d} + \frac{l_d}{2} \nabla d^s \cdot \nabla d^s$$

Compressive-shear crack density

The energy dissipated by cracks :

$$D(d^s, d^t) = \int_{\Omega} [g_c^t \gamma_d^t(d^t, \nabla d^t) + g_c^s \gamma_d^s(d^s, \nabla d^s)] dV$$

Regularized crack fields

Poroelastic model for the undamaged material (Coussy, 2010) :

$$d\boldsymbol{\sigma}^0 = d\boldsymbol{\sigma}^{b0} - bS_w dp_w I$$

$$dp_w = M_{ww} \left[-bS_w d\boldsymbol{\varepsilon}_v + \left(\frac{dm_w}{\rho_w} \right) \right]$$

➤ The capillary pressure ():

$$p_c = -p_w$$

➤ The extended Biot's effective stress :

$$d\boldsymbol{\sigma}^{b0} = d\boldsymbol{\sigma}^0 + bS_w dp_w I = \mathbb{C}^{b0} : d\boldsymbol{\varepsilon}$$

➤ The water saturation degree (van Genuchten, 1980) :

$$S_w = S_r + S_e(1 - S_r)$$

$$S_e = \left[1 + \left(\frac{p_c}{p_{cr}} \right)^n \right]^{-m}$$

The total energy functional of partially saturated cracked material

$$E(\boldsymbol{\varepsilon}, m_w, d^t, d^s) = \underbrace{\int_{\Omega} \psi(\boldsymbol{\varepsilon}, m_w, d^t, d^s) dV}_{\text{stored energy}} + \underbrace{\int_{\Omega} \mathcal{D}(d^t, d^s) dV}_{\text{cracking dissipation}}$$

$$\psi(\boldsymbol{\varepsilon}, m_w, d^t, d^s) = \underbrace{\psi_{eff}(\boldsymbol{\varepsilon}^e, d^t, d^s)}_{\text{Skeleton deformation}} + \underbrace{\psi_{fluids}(\boldsymbol{\varepsilon}^e, m_w, m_g)}_{\text{Fluid mass change}}$$

Stored energy for partially saturated media with cracks

- The stored elastic energy of porous medium:

$$\psi_{eff}(\boldsymbol{\varepsilon}, d^t, d^s) = g(d^t) \boxed{W_+^b(\boldsymbol{\varepsilon})} + g(d^s) \boxed{W_-^b(\boldsymbol{\varepsilon})}$$

$$W_+^b(\boldsymbol{\varepsilon}) = \frac{1}{2} \boldsymbol{\sigma}_+^b : \boldsymbol{\varepsilon}$$

Tensile crack driving energy

$$W_-^b(\boldsymbol{\varepsilon}) = \frac{1}{2} \boldsymbol{\sigma}_-^b : \boldsymbol{\varepsilon}$$

Shear crack driving energy

- The degradation function (Miehe et al. 2010) :

$$g(d^\alpha) = (1 - d^\alpha)^2$$

- The Decomposition of effective stress tensors :

$$\boldsymbol{\sigma}_\pm^b = \sum_{a=1}^3 \langle \sigma_a \rangle_\pm \mathbf{n}_a \otimes \mathbf{n}_a$$

- The energy due to fluid mass change :

$$\psi_{fluids}(\boldsymbol{\varepsilon}, m_w, \boldsymbol{d}^t, \boldsymbol{d}^s) \equiv \psi_{fluids}(\boldsymbol{\varepsilon}, m_w) = \frac{1}{2} M_{ww} \left[b S_w \boldsymbol{\varepsilon}_v - \left(\frac{m}{\rho} \right)_w \right]^2$$

$$\Pi(\dot{\mathbf{u}}, \dot{m}_w, \dot{m}_g, \dot{d}^t, \dot{d}^s) = \dot{E}(\dot{\mathbf{u}}, \dot{m}_w, \dot{m}_g, \dot{d}^t, \dot{d}^s) - \dot{P}_{ext}$$

Governing equations for phase-field variables

$$-g'_t(d^t) \boxed{W_+^b} - g_c^t \left[\frac{d^t}{l} - l \operatorname{div}(\nabla d^t) \right] = 0$$

$$\boxed{W_+^e = \frac{1}{2} \boldsymbol{\sigma}_+ : \boldsymbol{\varepsilon}^e}$$

$$\mathcal{H}^t = \max_{t \in [0, t]} W_+^e$$

$$-2(1 - d^t) \mathcal{H}^t - g_c^t \left[\frac{d^t}{l} - l \operatorname{div}(\nabla d^t) \right] = 0$$

$$-g'_s(d^s) \boxed{W_-^b} - g_c^s \left[\frac{d^s}{l} - l \operatorname{div}(\nabla d^s) \right] = 0$$

$$\boxed{W_-^s = \frac{1}{2G} \left\langle \frac{\sigma_3^- - \sigma_1^-}{2 \cos \varphi} + \frac{\sigma_3^- + \sigma_1^-}{2} \tan \varphi - c \right\rangle_+^2}$$

$$\mathcal{H}^s = \max_{t \in [0, t]} W_-^s$$

$$-2(1 - d^s) \mathcal{H}^s - g_c^s \left[\frac{d^s}{l} - l \operatorname{div}(\nabla d^s) \right] = 0$$

$$\Pi(\dot{\mathbf{u}}, \dot{m}_w, \dot{m}_g, \dot{d}^t, \dot{d}^s) = \dot{E}(\dot{\mathbf{u}}, \dot{m}_w, \dot{m}_g, \dot{d}^t, \dot{d}^s) - \dot{P}_{ext} = 0$$

Hydro-mechanics coupling functions for partially saturated medium

$$p_w - p_{w0} = M_w \left[-b_w \boldsymbol{\varepsilon} \mathbf{I} + \frac{m_w}{\rho_w} \right]$$



Darcy's law and mass conservation

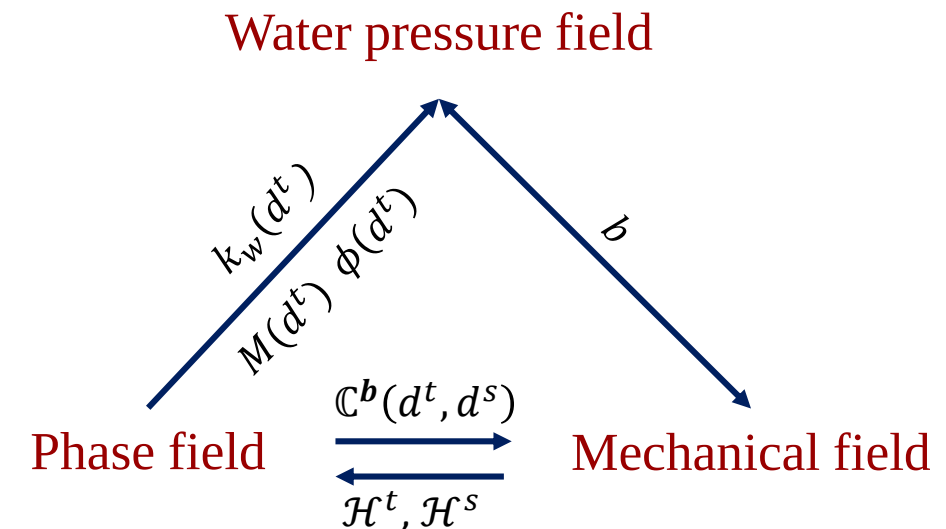
$$bS_w \dot{\varepsilon}_v + \frac{1}{M} \dot{p}_w = \frac{k_r k_w}{\mu_w} \cdot \text{div}(\nabla p_w - \rho_w \vec{g})$$

$$\text{div}(\boldsymbol{\sigma}) + \vec{f} = 0$$

$$\boldsymbol{\sigma} - \boldsymbol{\sigma}_0 = \mathbb{C}^b(d^t, d^s) : \boldsymbol{\varepsilon} - bS_w(p_w - p_{w0})\mathbf{I}$$

Effects of phase field on hydraulic parameters:

- Permeability: $k_w(d^t) = k_w^0 \exp(d^t)$
- Porosity: $\phi(d^t) = \phi^0 + (1 - \phi^0) d^t$
- Scalar parameter: $\frac{1}{M(d^t)} = \frac{s_l^2 [b - \phi(d^t)]}{K_s} + \frac{s_l \phi(d^t)}{K_f} - \phi(d^t) \frac{\partial s_l}{\partial p_c}$



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PART THREE

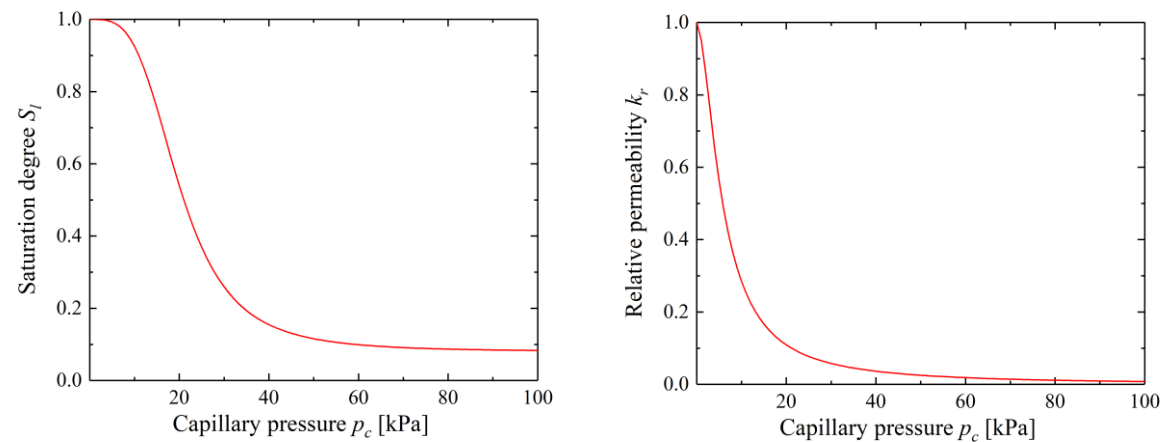
NUMERICAL MODELING

Analysis of rainfall induced landslides

Hydro-mechanical parameters:

λ	μ	K_w	ϕ	b	k_{pl}
2.9 GPa	0.7 GPa	2.2×10^9 Pa	0.38	1.0	5×10^{-12} m ²

Water retention and relative permeability curves



Phase-field parameters:

Critical energy g_c^t	Critical energy g_c^s	Crack length scale l
224 N/m	364 N/m	0.25m

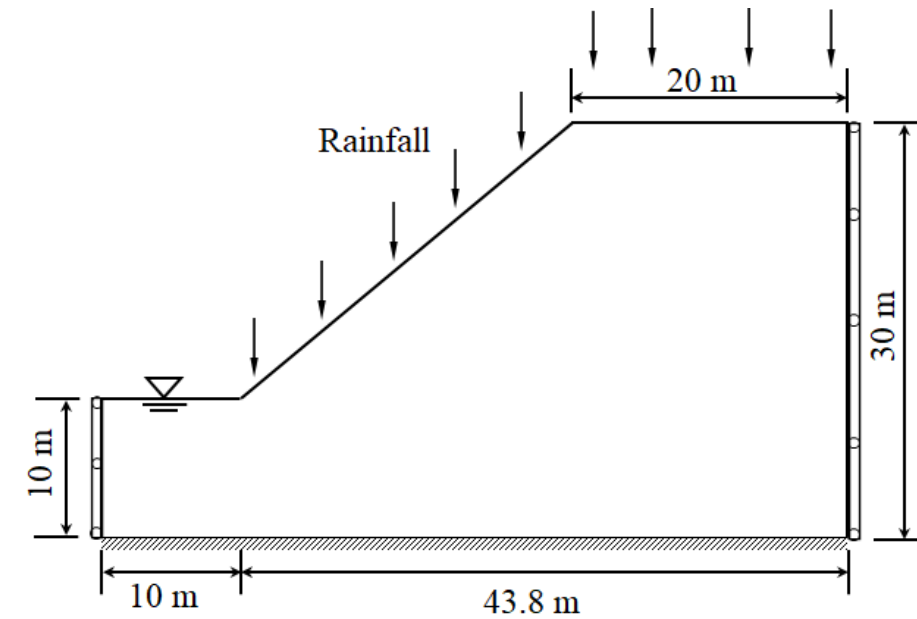


Figure - Boundary conditions

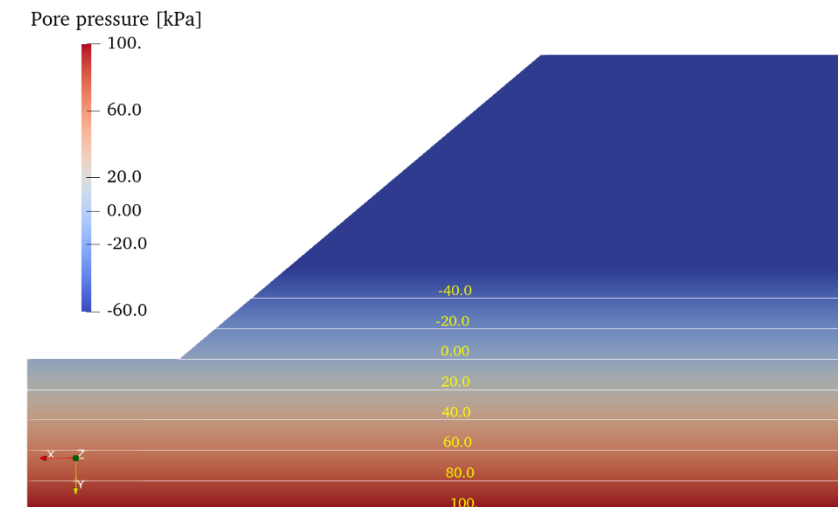


Figure - Initial distribution of pore pressure

Analysis of rainfall induced landslides

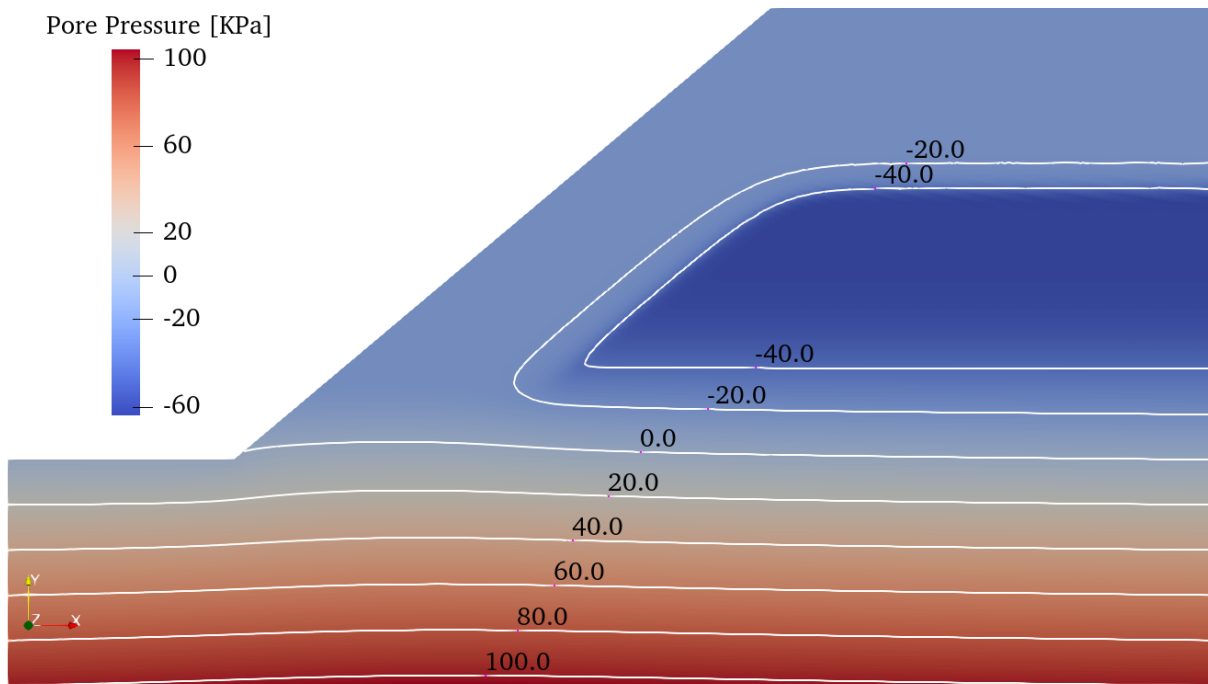


Figure - Distribution of pore pressure without damage (after 66h)

Rainfall infiltration:

- Increment of underground water table
- Partially saturated $\longrightarrow$ fully saturated (toe of the slope)

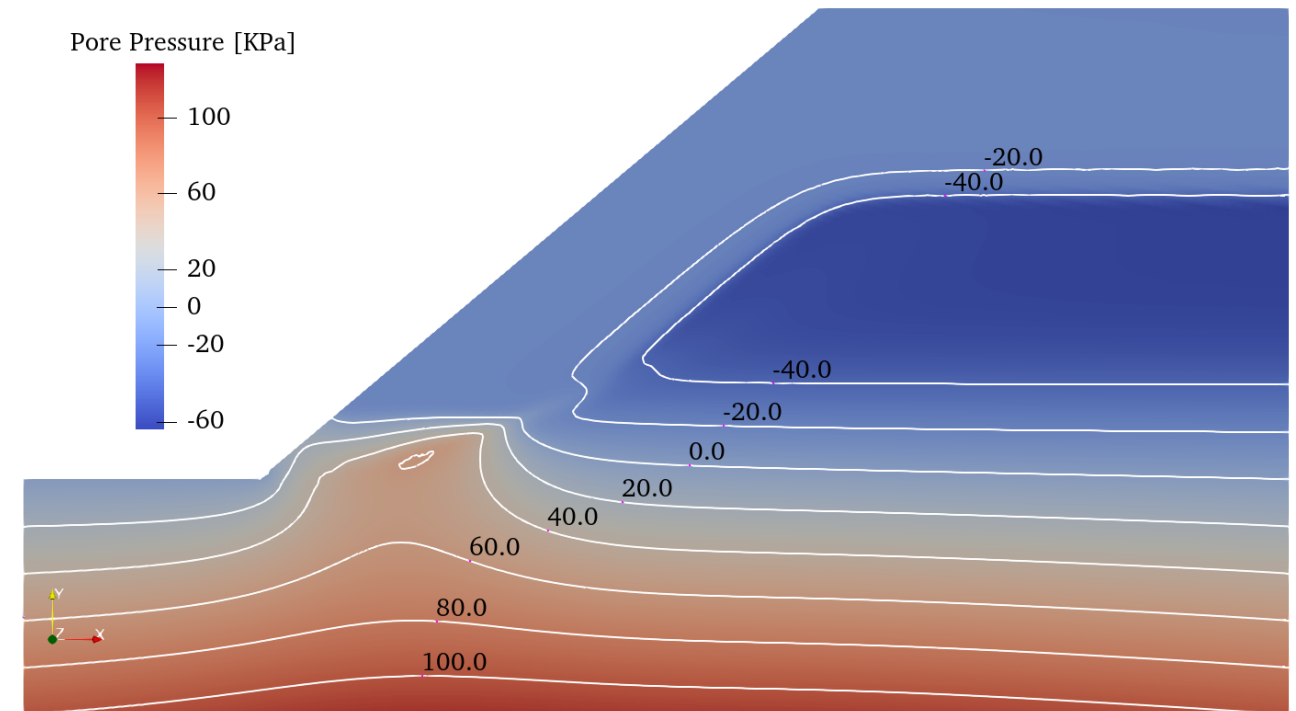


Figure - Distribution of pore pressure when slope failure occurs (after 65.5h)

Pore pressure $\longleftrightarrow$ cracks

Analysis of rainfall induced landslides

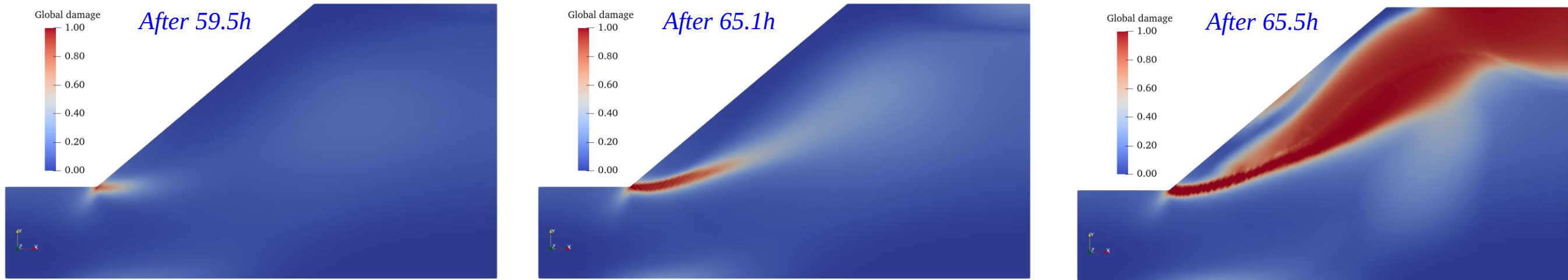


Figure - Distribution of global damage

- Onset of cracks: Around the toe of the slope
- Cracks path: Toe of slope → top of slope

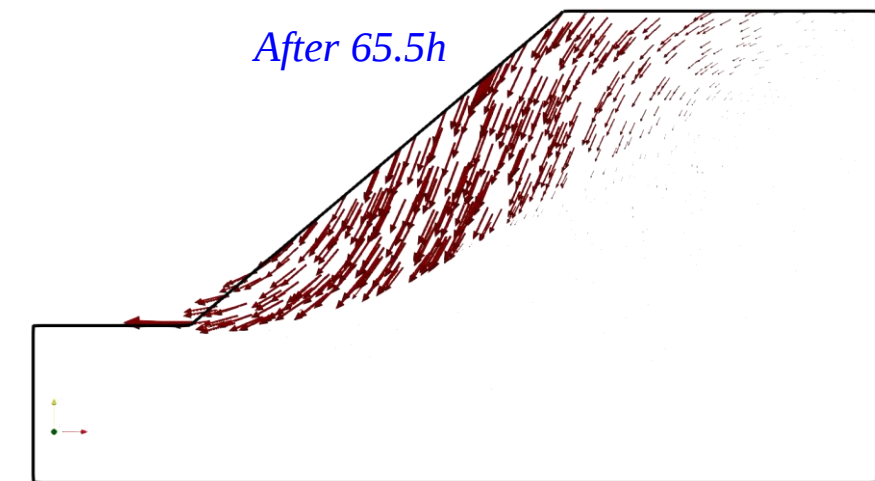
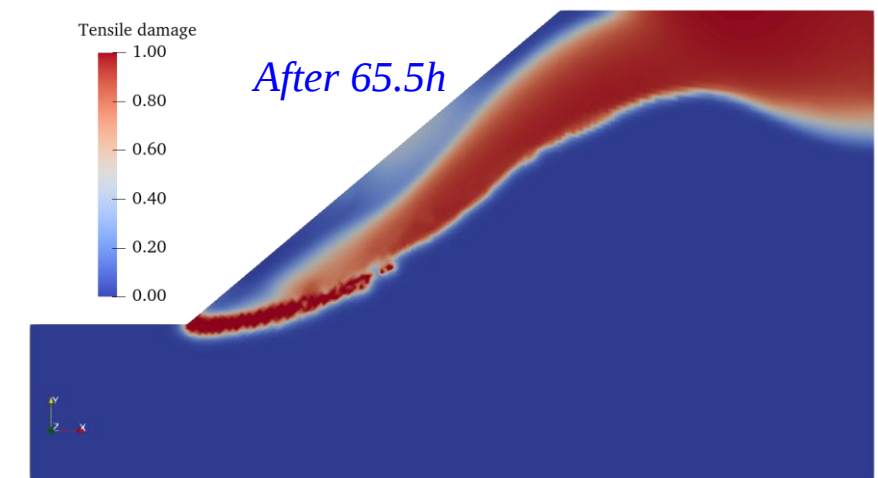
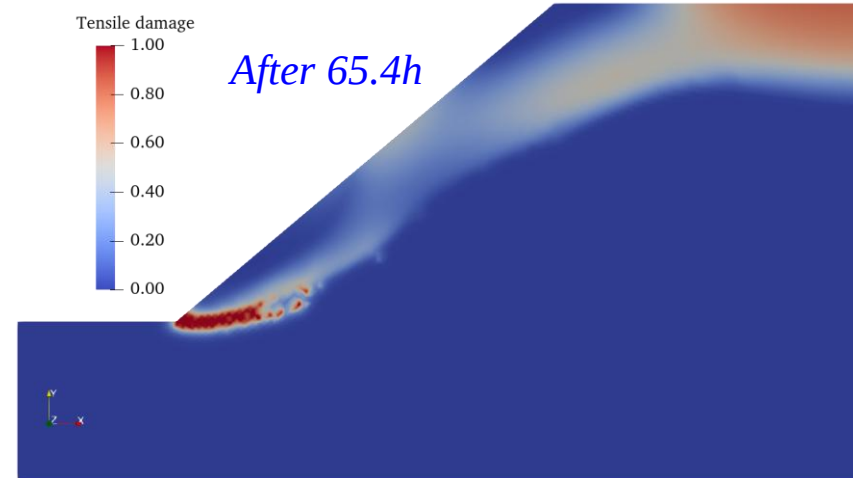
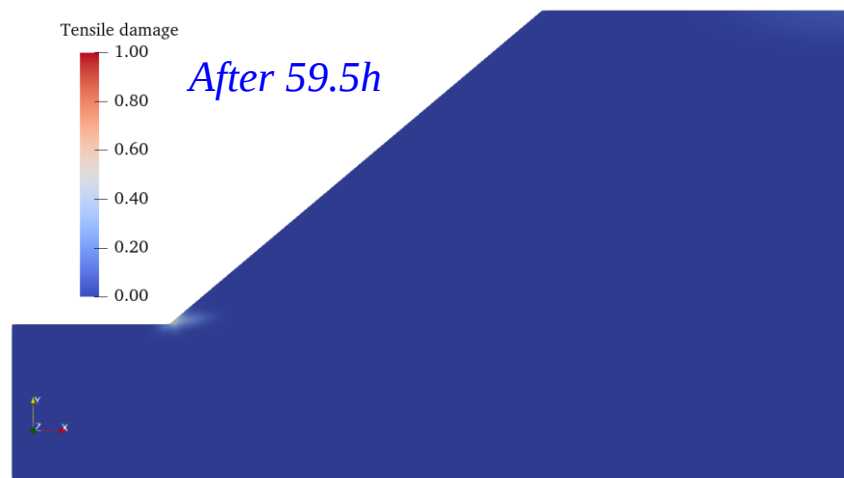
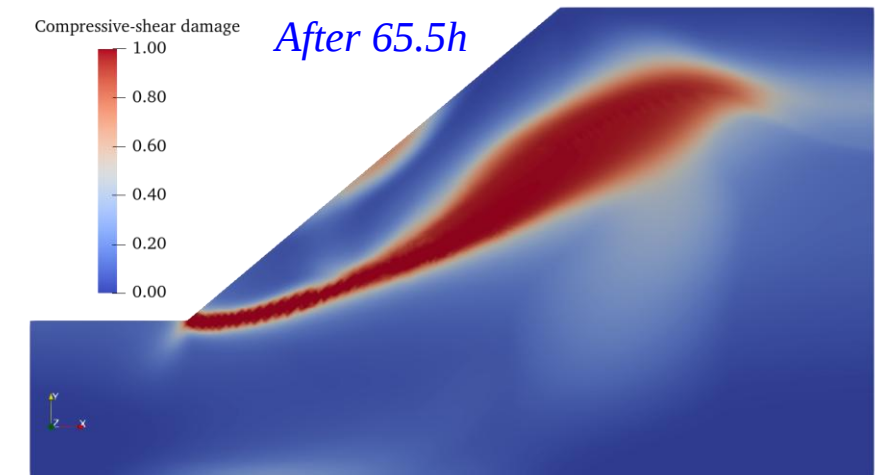
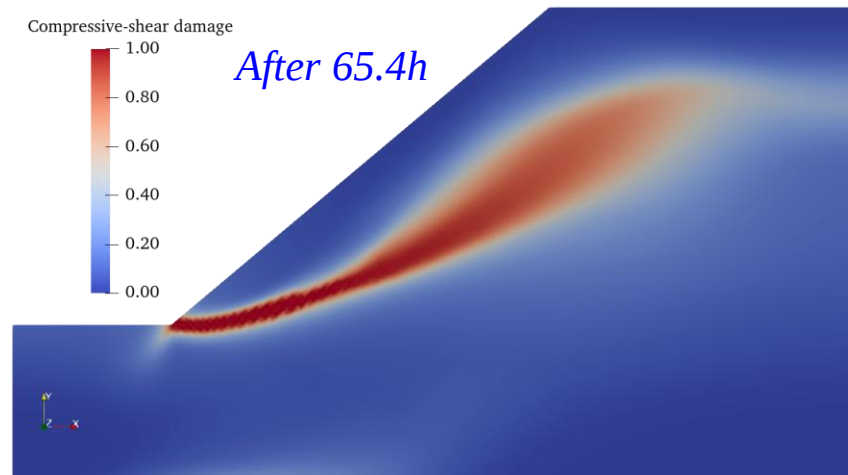
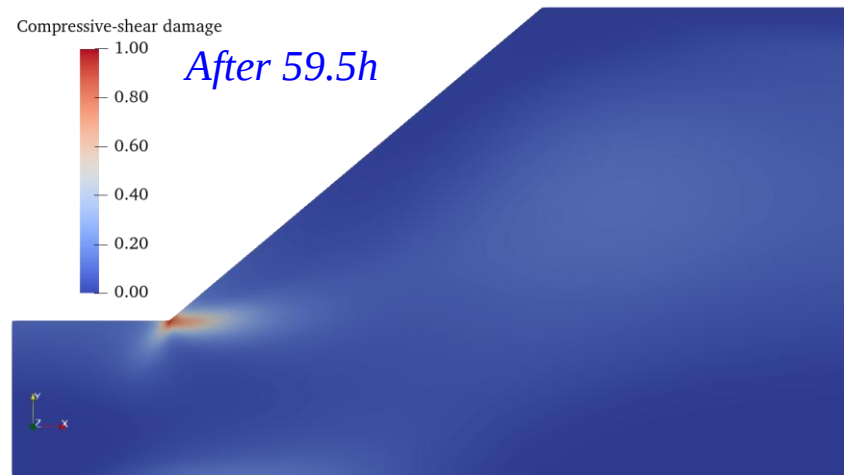


Figure - Displacement vector

Analysis of rainfall induced landslides

Compressive- shear cracks



Tensile cracks

Analysis of rainfall induced landslides

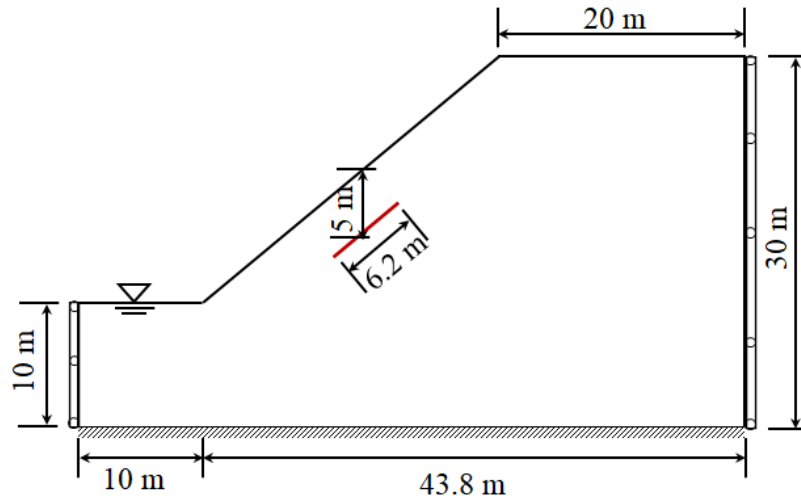


Figure – Distribution of pre-crack

Influences of pre-crack:

- Growth of cracks
- Two-step failure pattern

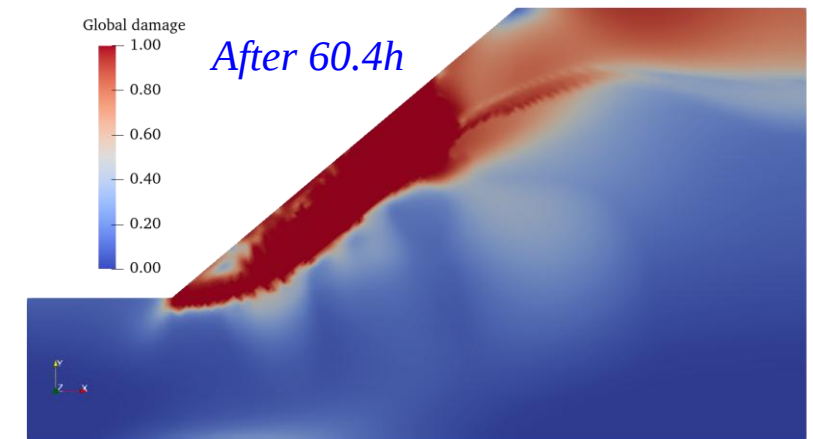
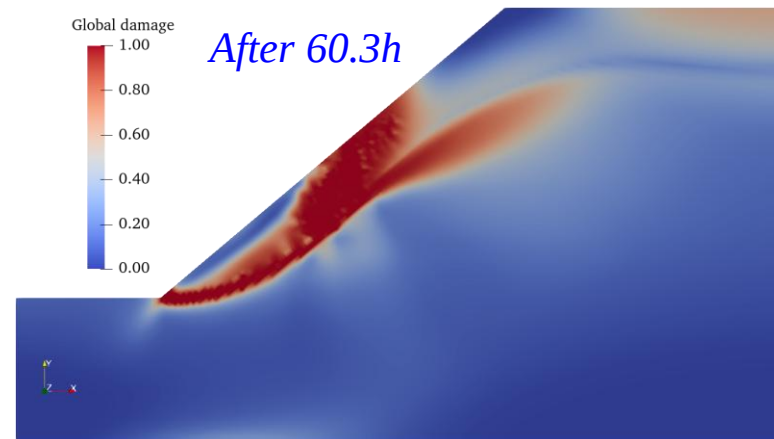
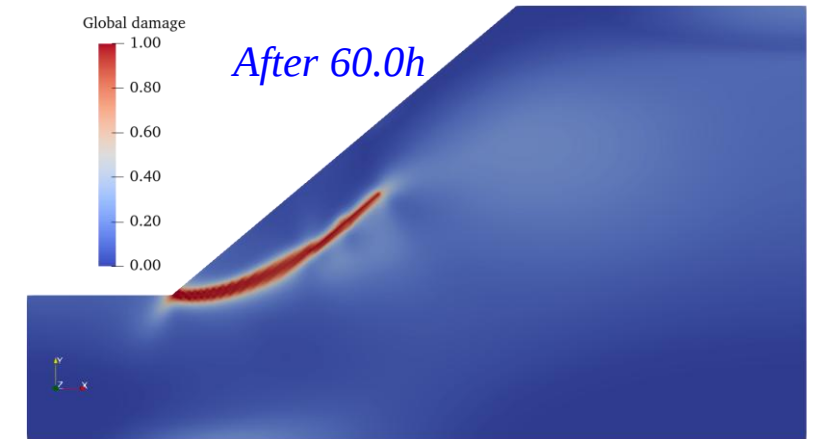
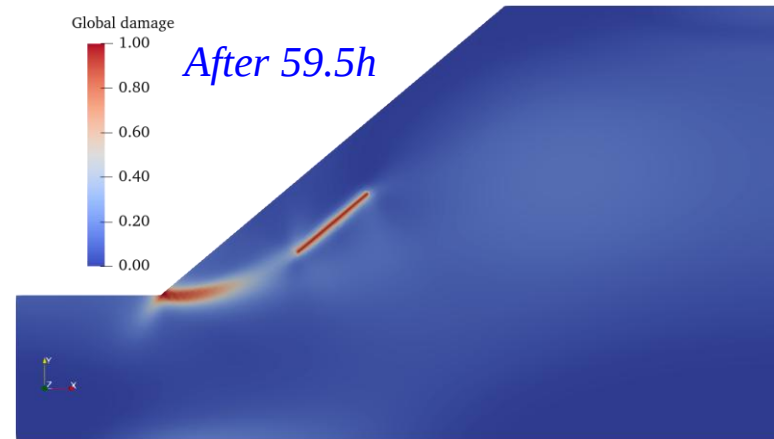
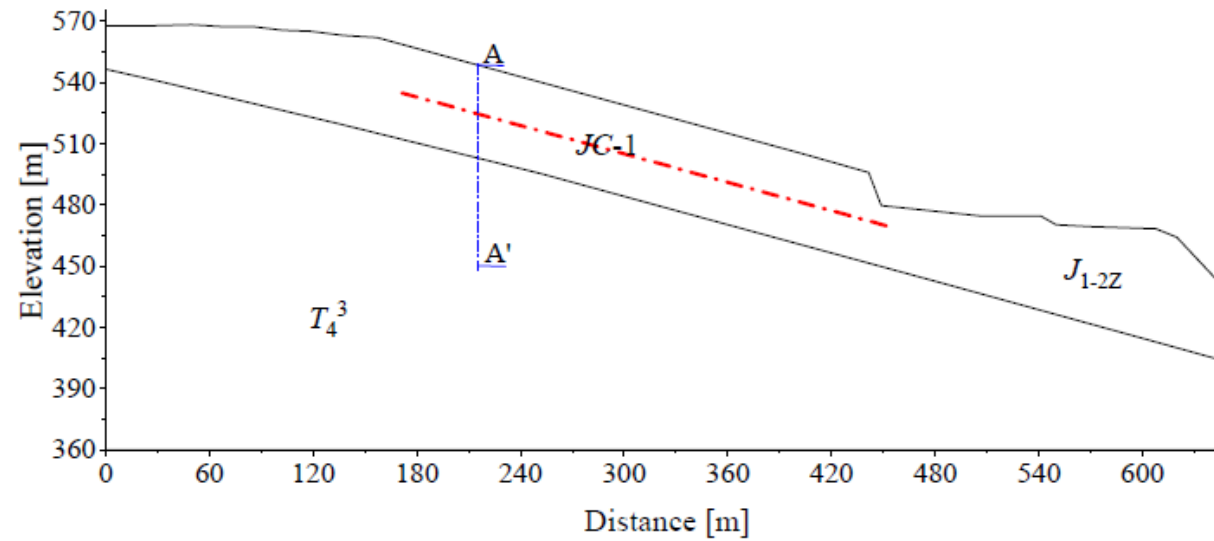


Figure - Distribution of global damage

Analysis of rainfall induced landslides

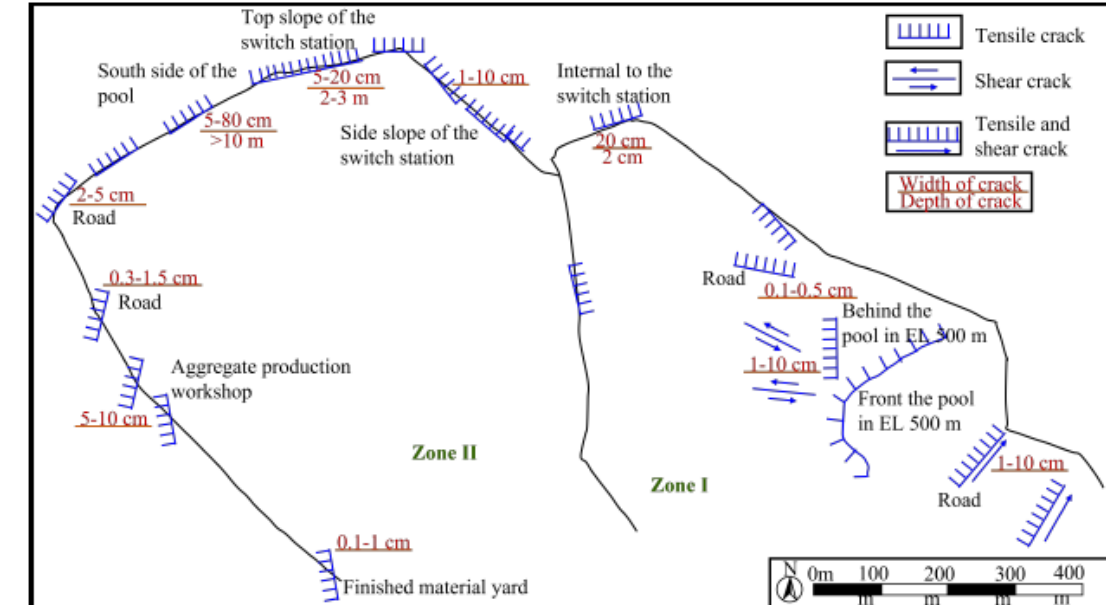
Description of numerical model:



$$\mathcal{H}_0^{t/s}(\mathbf{x}) = \begin{cases} \frac{g_c^{t/s}}{2l_d} \frac{d}{1-d} \left(1 - \frac{2L(\mathbf{x})}{l_d} \right) & L(\mathbf{x}) \leq \frac{l_d}{2} \\ 0 & \text{otherwise} \end{cases}$$

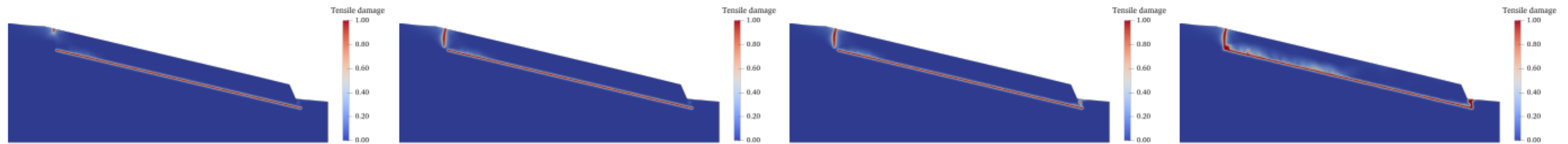
Field investigation of cracks distribution

(Zhang et al., 2018):



Analysis of rainfall induced landslides

Distribution of tensile damage



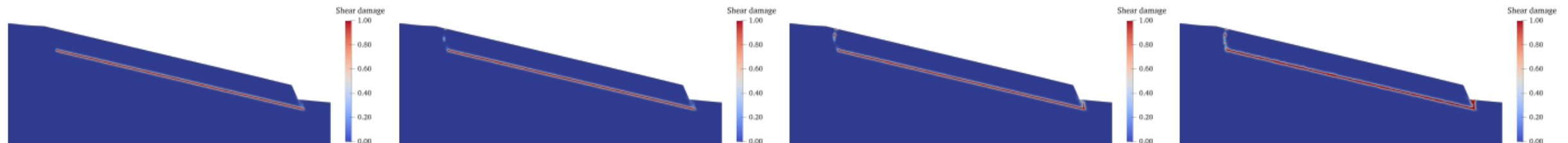
(a) $RF = 568.9\text{mm}$

(b) $RF = 598.7\text{mm}$

(c) $RF = 645.5\text{mm}$

(d) $RF = 651.3\text{mm}$

Distribution of shear damage



(a) $RF = 568.9\text{mm}$

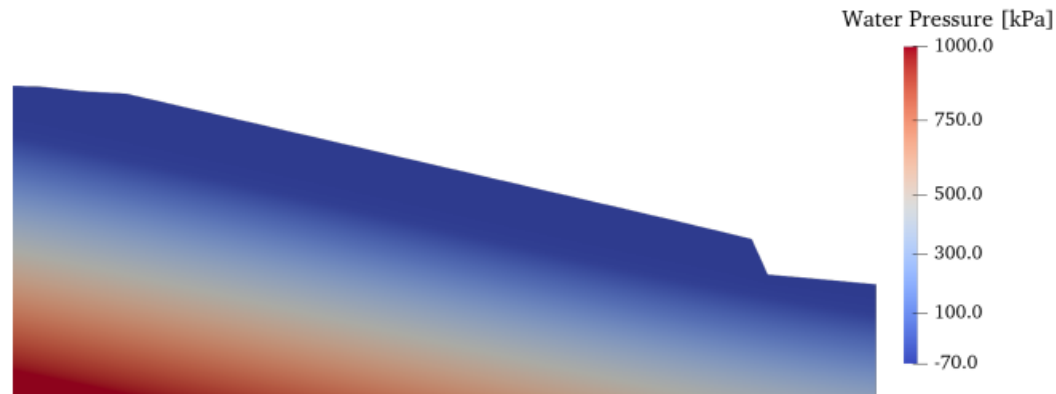
(b) $RF = 598.7\text{mm}$

(c) $RF = 645.5\text{mm}$

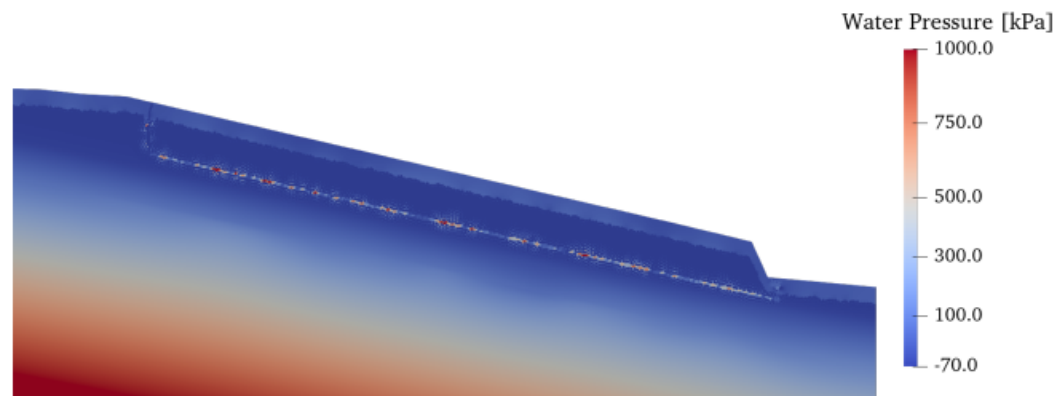
(d) $RF = 651.3\text{mm}$

Analysis of rainfall induced landslides

Distribution of pore water pressure:

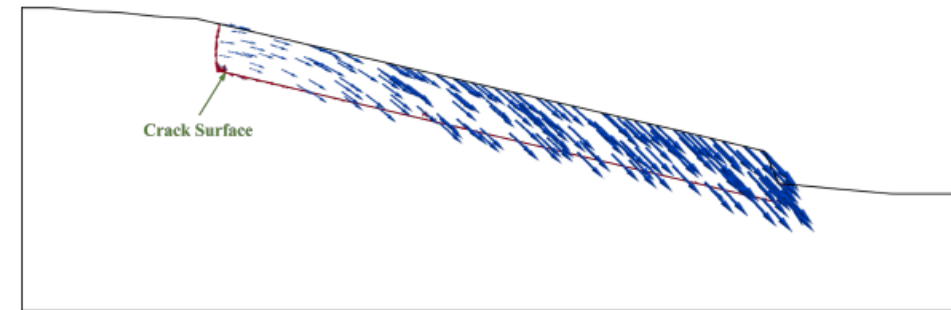


(a) Initial state

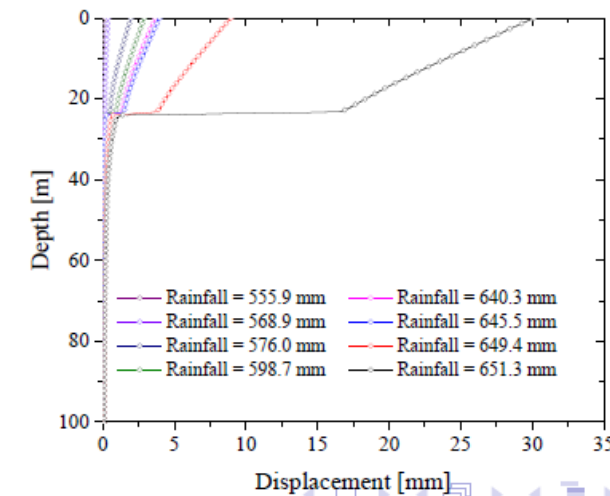


(b) RF = 651.3mm

Distribution of displacement:



(a) Displacement vector at RF = 651.3mm



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PART FOUR

CONCLUSIONS AND PERSPECTIVES

Conclusions

- The proposed method is able to describe the initiation and propagation of localized damage zones and cracks due to rainfall.
- It was found that the shear cracking was the principal failure mechanism of landslides.
- The existence of initial weak zones and fractures enhances the failure process and also affects the cracking pattern

Perspectives

- Application the proposed numerical method into analysis of reality landslides;
- Considering the material in a slope as a heterogeneous material;
- Proposing a time-dependent phase-field method to simulate the long-term behavior of the slope;
- Taking into account hydrodynamic effects

Thank you for your attendance !

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